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RMetS

Impact of non-stationarity on hybrid ensemble filters: A study with a doubly stochastic advection-diffusion-decay model

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Abstract

Effects of non-stationarity on the performance of hybrid ensemble filters is studied. (By hybrid filters we mean those which blend ensemble covariances with some other regularizing covariances.) To isolate effects of non-stationarity from effects due to nonlinearity (and the non-Gaussianity it causes), a new doubly stochastic advection-diffusion-decay model (DSADM) is proposed. The model is hierarchical: it is a linear stochastic partial differential equation whose coefficients are random fields defined through their own stochastic partial differential equations. DSADM generates conditionally Gaussian spatiotemporal random fields with a tunable degree of non-stationarity in space and time. DSADM allows the use of the exact Kalman filter as a baseline benchmark.

In numerical experiments with DSADM as the "model of truth", the relative importance of the three kinds of covariance blending is studied: with static, time-smoothed, and space-smoothed covariances. It is shown that the stronger the non-stationarity, the less useful the static covariance matrix becomes and the more beneficial the time-smoothed covariances are. Time-smoothing of background-error covariances proved to be systematically more useful than their space-smoothing. Under non-stationarity, a filter that extends the (previously proposed by the authors) Hierarchical Bayes Ensemble Filter and accommodates the three covariance-blending techniques is shown to outperform all other configurations of the filters tested.

KEYWORDS

advection-diffusion-decay model, data assimilation, EnKF, hierarchical modelling, hybrid ensemble filters, hierarchical Bayes ensemble filter, non-stationary spatiotemporal random field

1 INTRODUCTION

Ensemble Kalman filters (EnKF) estimate background-error covariances from an ensemble of forecasts. The resulting sample covariances are noisy and rank deficient if the ensemble size is not greater than the dimensionality of the covariance matrix (which is typically the case in geosciences). A kind of regularization (that is, introduction of additional information to stabilize the solution) is needed to make them a useful estimate of true background-error covariances. Two

regularization techniques are most popular: localization and hybridization. Our focus in this study is on the second technique, in which sample covariances are blended (mixed) with some regularizing covariances and which we therefore call "covariance blending."

From the literature, we can identify three main types of covariance blending. First, hybrid ensemble-variational filters (EnVar) employ blending with static "climatological" covariances (Hamill and Snyder, 2000; Lorenc, 2003; Buehner *et al.*, 2013). In statistical literature this regularization device

is known as a shrinkage estimator (Ledoit and Wolf, 2004). Second, in the Hierarchical Bayes Ensemble Filter (HBEF) by Tsyrulnikov and Rakitko (2017), ensemble covariances are blended with time-smoothed recent-past covariances. The idea of accommodating recent-past background-error statistics has also been explored by other authors: Gustafsson *et al.* (2014) used time-lagged ensemble members, Berre *et al.* (2015) used ensemble members from the past 4 days to increase the ensemble size, Bonavita *et al.* (2016) used ensemble covariances from the previous 12 days to estimate their parametric covariance model, and Lorenc (2017) found that using time-lagged and time-shifted perturbations increases the effective ensemble size. Third, Buehner and Charron (2007), Ménétrier *et al.* (2015), Ménétrier and Auligné (2015) studied spatial smoothing of ensemble covariances, which also can be considered as covariance blending, this time with neighbouring covariances in space.

With any of the above kinds of covariance blending, the intention is to reduce sampling noise and increase the rank of the covariance matrix – but at the expense of reducing flow dependence (the main advantage of ensemble statistics) in the blended covariances. The question arises: what are the optimal weights of the regularizing covariances in the blend and which factors do they depend on?

First of all, we notice that if a data assimilation system is stationary, that is, its true (actual) spatial error covariances do not change in time, then there is no need for an (on-line) ensemble at all. Indeed, with constant-in-time background-error covariances, all one has to do is to carefully estimate (off-line) their spatial covariance matrix (through temporal averaging of innovation-based or ensemble covariances). In the Gaussian case, the estimated static covariance matrix will be optimal (e.g. Theorem 2.38 in Kumar and Varaiya, 2015, gives a formal statement). Thus, we conclude that only under *non-stationarity* can ensemble filters outperform data assimilation schemes that employ static background-error covariances. True background-error covariances in a data assimilation scheme are non-stationary either if the system is governed by a non-stationary model or the observation network is changing in time (or both, of course). In this research, our focus is on non-stationarity caused by the model. To study the impact of non-stationarity on the optimal design of hybrid filters, we seek a toy model that is capable of producing spatiotemporal fields with tunable non-stationarity.

Most toy models used in data assimilation studies are nonlinear. The simplest models have just a few variables, like the Ikeda model (Hansen and Smith, 2001), the double well model (Miller *et al.*, 1994), and the most popular three-variable Lorenz'63 model (Lorenz, 1963). Models defined on a 1D spatial domain include, among others, the viscous Burgers equation (Apte *et al.*, 2010), the Korteweg–de Vries equation (Muccino and Bennett, 2002), and the most popular Lorenz'96 and Lorenz'2005 models (Lorenz and Emanuel, 1998; Lorenz, 2005). In nonlinear models, the tangent-linear model operator, generally, varies in time, giving rise to non-stationarity (flow-dependence) of filtering probability distributions. However, with nonlinear models, non-stationarity is inevitably intertwined with nonlinearity of the evolution of filtering errors and their ensuing non-Gaussianity. To disentangle effects due to nonlinearity from effects of non-stationarity, we propose a model that gives rise to non-stationary filtering distributions while being linear and Gaussian.

We build the model on the time-discrete *scalar* (i.e. one-variable) model introduced by Tsyrulnikov and Rakitko (2017):

$$
\xi_k = F_k \xi_{k-1} + \sigma_k \alpha_k,\tag{1}
$$

where ξ_k is the (scalar) true system state, k labels the time instant, F_k is the (scalar) model operator, σ_k is the standard deviation of the forcing, and $\alpha_k \sim \mathcal{N}(0, 1)$ is the driving white noise. The model is *doubly stochastic* (e.g. Tjøstheim, 1986), which means that not only the forcing α_k is random, the coefficients F_k and σ_k are random sequences by themselves, each defined through its own stochastic model similar to Equation (1) but with a constant model operator and constant magnitude of the forcing. *Conditionally on the processes* F_k and σ_k (i.e. after their realizations are computed and kept fixed), Equation (1) is a linear model with time-varying coefficients (e.g. Chatfield, 2016, section 11.2.1), whose solution is a non-stationary random process. In this study we extend the model Equation (1) to the spatiotemporal context. The result is a new doubly stochastic advection-diffusion-decay model (DSADM), whose solutions are non-stationary conditionally Gaussian random fields.

The general idea of non-stationary random field modelling by assuming that parameters of a spatial or spatiotemporal model are random fields themselves is discussed by several authors. Piterbarg and Ostrovskii (2013) study an advection-diffusion model whose coefficients are random fields with prescribed covariance functions. Lindgren *et al.* (2011) allow a length-scale parameter and a variance parameter of their stochastic partial differential equation to slowly vary in space (by expanding them in a set of basis functions). Banerjee *et al.* (2015, section 11.6) mention the possibility of formulating a stochastic differential equation for parameters of another stochastic differential equation. Roininen *et al.* (2019) use a stochastic elliptic equation to model a spatial random field and introduce a hypermodel for its local length-scale to achieve non-stationarity; the hypermodel employs the same stochastic elliptic equation. Dunlop *et al.* (2018) discuss (among others) more general Gaussian hierarchical processes, which are defined through a hierarchy of levels so that at each level, the conditional probability distribution is Gaussian given the previous (parent) level. Our innovation is the non-stationary hierarchical model with spatiotemporal stochastic partial differential equations at two levels in the hierarchy. The particular pattern of the non-stationarity (i.e. how the field's variance, local spaceand time-scales, etc. vary in space-time) is random, and its spatiotemporal structure is highly tunable by the model's hyperparameters.

The rest of the paper is organized as follows. In Sections 2–5 we motivate, describe, and examine in numerical experiments the new DSADM model. In Section 6 we introduce a new hybrid-HBEF (named HHBEF) filter, which accommodates the three covariance-blending techniques mentioned above. In Section 7 we present results of numerical experiments with hybrid ensemble filters showing the crucial impact of non-stationarity on the filters' performance and optimal design. Efficient numerical implementation of hybrid filters is beyond the scope of this study.

The R code of the doubly stochastic model, the filters, and the R scripts that produced this paper's numerical results are available from [https://github.com/cyrulnic/](https://github.com/cyrulnic/NoStRa) [NoStRa](https://github.com/cyrulnic/NoStRa) (accessed 8 May 2019). More detailed texts on the DSADM's numerical scheme, generation of initial conditions, specification of model parameters, etc. can also be found there.

2 STOCHASTIC ADVECTION-DIFFUSION-DECAY MODEL WITH CONSTANT COEFFICIENTS

To introduce notation and motivate the new doubly stochastic advection-diffusion-decay model, we theoretically examine here statistics of solutions to a *stationary* stochastic advection-diffusion-decay model with non-stochastic and constant coefficients (parameters). Specifically, we identify how the model parameters affect the variance, the spatial scale, and the time-scale of the solution. This is a preparatory section with largely background material.

2.1 Model

The model is the following stochastic partial differential equation (Whittle, 1986, chapter 20, section 3; Lindgren *et al.*, 2011, equation 17; Sigrist *et al.*, 2015):

$$
\frac{\partial \xi}{\partial t} + U \frac{\partial \xi}{\partial s} + \rho \xi - v \frac{\partial^2 \xi}{\partial s^2} = \sigma \alpha, \tag{2}
$$

where *t* is time, *s* is the spatial coordinate on the circle $\mathbb{S}^1(R)$ of radius R, U is the advection velocity, ρ is the decay (damping) parameter, v is the diffusion parameter, $\alpha(t, s)$ is the standard white (in space and time) Gaussian noise, and σ is the intensity of the forcing. The four parameters $\theta = (U, \rho, \nu, \sigma)$ are constant in space and time.

2.2 Stationary spatiotemporal statistics

We start with rewriting Equation (2) using the material time derivative (i.e., along the characteristic $s = s_0 + Ut$), or, equivalently, switching to the Lagrangian frame of reference by making the change of variables $(t, s) \mapsto (t, s - Ut)$:

$$
\frac{\mathrm{d}\xi}{\mathrm{d}t} + \rho\xi - \nu \frac{\partial^2 \xi}{\partial s^2} = \sigma\alpha. \tag{3}
$$

Next, we employ the spectral expansion in space (retaining only those spectral components that are resolved on the regular spatial grid with n points),

$$
\xi(t,s) = \sum_{m=-n/2+1}^{n/2} \widetilde{\xi}_m(t) e^{ims/R}
$$
 (4)

and

$$
\alpha(t,s) = \sum_{m=-n/2+1}^{n/2} \widetilde{\alpha}_m(t) e^{ims/R},
$$
\n(5)

where i is the imaginary unit and $\tilde{\xi}_m(t)$ and $\tilde{\alpha}_m(t)$ are the (complex) spectral coefficients. It can be shown (e.g. Tsyrulnikov and Gayfulin, 2016, appendix A.4) that $\tilde{\alpha}_m(t)$ are independent standard complex-white-noise processes $\omega_m(t)$ with the common intensity $a = 1/\sqrt{2\pi R}$:

$$
\widetilde{\alpha}_m(t) = a \omega_m(t). \tag{6}
$$

Now, we substitute Equations (4)–(6) into Equation (3), getting

$$
\frac{\mathrm{d}\tilde{\xi}_m}{\mathrm{d}t} + \left(\rho + \frac{\nu}{R^2}m^2\right)\tilde{\xi}_m = a\sigma\omega_m(t). \tag{7}
$$

This is the spectral-space form of the model Equation (3). It is easily seen that if $\rho + (v/R^2)m^2 > 0$, the solutions to Equation (7) for different m become, after an initial transient, mutually independent stationary zero-mean random processes.[1](#page-2-0)

This implies that the physical-space solution $\xi(t, s)$ becomes a zero-mean random field that is stationary in time and space. Note that by definition, the zero-mean real-valued random field (process) $\xi(t, s)$ is (second-order) stationary if its spatiotemporal covariance function

$$
\gamma(t_1, s_1; t_2, s_2) = \mathbb{E}\,\xi(t_1, s_1)\,\xi(t_2, s_2)
$$

¹ Indeed, the influence of the initial condition on $\tilde{\xi}_m(t)$ exponentially decays in time, leaving $\tilde{\xi}_m(t)$ dependent only on the noise process $\omega_m(t')$ for $0 \le t' \le t$. The mutual independence of $\tilde{\xi}_m(t)$ for different *m* then follows from the mutual independence of the driving noises $\omega_m(t)$.

is invariant under translations:

$$
\mathbb{E}\,\xi(t_1,s_1)\,\xi(t_2,s_2)=\mathbb{E}\,\xi(t_1+u,s_1+v)\,\xi(t_2+u,s_2+v)
$$

and thus is a function of the space and time shifts only:

$$
\gamma(t,s) = \mathbb{E}\,\xi(t_1,s_1)\,\xi(t_1+t,s_1+s).
$$

Here periodicity in the spatial coordinate s is of course assumed, $\xi(t, s) = \xi(t, s + 2\pi R)$.

Each elementary stochastic process $\tilde{\xi}_m(t)$ is an Ornstein–Uhlenbeck process (e.g. Arnold, 1974, section 8.3) with the stationary covariance function

$$
B_m(t) = b_m \cdot e^{-|t|/\tau_m},\tag{8}
$$

where the spectral variances b_m are

$$
b_m = \frac{a^2 \sigma^2}{2} \cdot \frac{1}{\rho + \frac{v}{R^2} m^2} \tag{9}
$$

and the spectral time-scales τ_m are

$$
\tau_m = \frac{1}{\rho + \frac{v}{R^2} m^2}.
$$
\n(10)

Note that Equation (10) provides the motivation for the inclusion of the decay term in the model. Indeed, with $\rho = 0$, the time-scale τ_0 would be infinitely large, which, on the finite sphere, is unphysical. (Specifying $b_0 = 0$ could resolve the problem but only at the expense of nullifying $\tilde{\xi}_0$, which is also unphysical.)

The stationary in space-time covariance function $\gamma(t, s)$ of the random field $\xi(t, s)$ can be easily derived from Equation (4) while utilizing the independence of the spectral processes $\tilde{\xi}_m(t)$, the spectral-space temporal covariance functions given in Equation (8), and returning to the Eulerian frame of reference:

$$
\gamma(t,s) = \sum_{m=-n/2+1}^{n/2} b_m e^{-|t|/\tau_m} e^{im(s-Ut)/R}.
$$
 (11)

Note that Equation (11) implies that the space-time correlations are *non-separable*, that is, they cannot be represented as a product of purely spatial and purely temporal correlations. Moreover, according to Equation (10), smaller spatial scales (i.e. larger wavenumbers m) correspond to smaller time-scales τ_m . This feature of space-time correlations ("proportionality of scales") is physically reasonable – as opposed to the simplistic and unrealistic separability of space-time correlations – and widespread in the real world (Tsyroulnikov, 2001, Tsyrulnikov and Gayfulin, 2017, and references therein).

Finally, from Equation (11), the stationary (steady-state) variance of $\xi(t, s)$ is

$$
\begin{aligned} \text{Var}\,\xi &\equiv \{\text{SD}(\xi)\}^2 = \sum_{m=-n/2+1}^{n/2} b_m \\ &\quad = \frac{a^2 \sigma^2}{2} \sum_{m=-n/2+1}^{n/2} \left\{ \rho + \frac{\nu}{R^2} m^2 \right\}^{-1}, \quad (12) \end{aligned}
$$

where SD stands for standard deviation.

2.3 Roles of model parameters

Firstly, we note that U does not impact the variance spectrum and the spectral time-scales of ξ (Equations (9) and (10)), its role is just to rotate the solution with the constant angular velocity U/R .

Secondly, Equation (9) implies that the *shape* of the spatial spectrum is

$$
b_m \propto \frac{1}{\rho + \frac{v}{R^2} m^2} \propto \frac{1}{1 + \left(\frac{m}{m_0}\right)^2},\tag{13}
$$

where $m_0 = R\sqrt{\rho/\nu}$ is the characteristic non-dimensional wavenumber, which defines the *width* of the spectrum and thus the field's length-scale. Therefore, the latter can be defined^{[2](#page-3-0)} as the inverse dimensional wavenumber m_0/R :

$$
L = \sqrt{\frac{v}{\rho}}.\tag{14}
$$

Thus, the ratio v/ρ controls the length-scale L. In addition, v/ρ impacts the temporal correlations. Indeed, a higher L implies a redistribution of the variance towards larger spatial scales (i.e. lower wavenumbers m). But as we noted, in the model Equation (2), larger spatial scales correspond to larger time-scales τ_m . As a result, a higher v/ρ leads to a larger time-scale as well as the length-scale L .

Thirdly, using Equation (14), we can rewrite Equation (10) as

$$
\tau_m = \frac{1}{\rho} \cdot \frac{1}{1 + \left(\frac{L_m}{R}\right)^2} \tag{15}
$$

This equation implies that, with L being fixed, all spectral time-scales τ_m are inversely proportional to ρ , which thus determines the physical-space Lagrangian time-scale T of the

² One can show that, for a dense enough spatial grid, the thus defined length-scale L almost coincides with the macroscale Λ_{ξ} defined below in Equation (30).

spatiotemporal random field ξ . We define T as the macroscale (e.g. Yaglom, 1987, equation 2.88) along the characteristic,

$$
T = \frac{1}{2\text{Var}\,\xi} \int_{-\infty}^{\infty} \gamma(t, Ut) \, \mathrm{d}t = \frac{\sum b_m \tau_m}{\sum b_m} = \frac{1}{\rho} \frac{\sum \left[1 + \left(\frac{Lm}{R}\right)^2\right]^{-2}}{\sum \left[1 + \left(\frac{Lm}{R}\right)^2\right]^{-1}},\tag{16}
$$

where the second equality is due to Equation (11), the third equality is due to Equations (9) and (15), and the summations are over *m* from $-(n/2) + 1$ to $n/2$.

Technically, Equations (14), (16), and (12) allow us to compute the *internal* model parameters ρ , v , and σ from the *externally* specified parameters L, T , and $SD(\xi)$. Conceptually, we summarize the above conclusions as follows.

- *U* does not affect the Lagrangian spatiotemporal covariances. It tilts the Eulerian spatiotemporal correlations towards the direction $ds = U dt$ in space-time.
- The ratio v/ρ determines the spatial scale L and impacts the time-scale T .
- With the ratio ν/ρ being fixed, ρ controls the time-scale T .
- With ρ and ν being fixed, σ determines the resulting process variance Var ξ .

These conclusions give us an idea how *local* properties of the spatiotemporal field statistics are going to be impacted if the parameters $\theta = (U, \rho, v, \sigma)$ become variable in space and time.

3 DOUBLY STOCHASTIC ADVECTION-DIFFUSION-DECAY MODEL (DSADM)

Here, we allow the parameters $\theta = (U, \rho, v, \sigma)$ of the model Equation (2) to be spatiotemporal random fields by themselves. The resulting model thus becomes a three-level hierarchical model (Wikle *et al.*, 1998; Banerjee *et al.*, 2015). At the first level is the random field in question $\xi(t, s)$ modelled conditionally on the second-level fields $\theta(t, s) = \{U(t, s), \rho(t, s), v(t, s), \sigma(t, s)\}\$, which are controlled by the third-level hyperparameters ϕ . So, to compute a realization of the pseudo-random field $\xi(t, s)$, we first specify the hyperparameters ϕ . Then, we compute realizations of the second-level (secondary or parameter) fields $\theta(t, s)$. Finally, we substitute the secondary fields for the respective parameters in Equation (2) and solve the resulting equation for the primary field $\xi(t, s)$.

The idea behind this extension of the basic model Equation (2) is the following. If the secondary fields $\theta(t, s)$ vary *smoothly* in space and time, then, locally, in a vicinity of some point in space-time (t_0, s_0) , the statistics of the field $\xi(t, s)$ will resemble that for the stationary model Equation (2) with constant parameters equal to $\theta(t, s)$ *frozen* at the point (t_0, s_0) (also Lindgren *et al.*, 2011, section 3.2). As the statistics of the model Equation (2) with constant parameters do depend on the parameters θ (Section 2), the resulting solution $\xi(t, s)$ to the model Equation (2) with variable parameters becomes non-stationary in space-time, with the degree of the non-stationarity controlled by the variability in the secondary fields $\theta(t, s)$.

3.1 First level of the hierarchy: the field in question

At the first level, $\xi(t, s)$ satisfies the basic Equation (2) with variable in space and time coefficients,

$$
\frac{\partial \xi}{\partial t} + U(t, s) \frac{\partial \xi}{\partial s} + \rho(t, s) \xi - v(t, s) \frac{\partial^2 \xi}{\partial s^2} = \sigma(t, s) \alpha(t, s). \tag{17}
$$

3.2 Second level of the hierarchy: parameter fields

Each secondary field $\theta(t, s)$ (i.e. one of the coefficients $U(t, s), \rho(t, s), \nu(t, s), \sigma(t, s)$ of the first-level Equation (17)) is modelled as the *transformed Gaussian field*: $\theta = g_{\theta}(\theta^*, \phi)$. Here g_{θ} is the (secondary-field specific) transformation function and $\theta^*(t, s)$ is the pre-transform Gaussian random field satisfying its own stochastic advection-diffusion-decay model Equation (2) with constant and non-random coefficients, U_{θ} , ρ_{θ} , v_{θ} , σ_{θ} (which are hyperparameters).

The pointwise transformation g_{θ} : $\theta^*(t, s) \mapsto \theta(t, s)$ is specified to be linear for U and nonlinear for the other three parameter fields. The transformation function involves additional hyperparameters: the "unperturbed" value of θ (a scalar) denoted by the overbar, $\overline{\theta}$ (such that $\theta(t, s) = \overline{\theta}$ if Var $\theta^* = 0$, and a few additional hyperparameters as described below.

Since the pre-transform fields $\theta^*(t, s)$ are governed by the models with *constant* coefficients, $\theta^*(t, s)$ are stationary in space-time. The transforms g_{θ} are defined to be independent of (t, s) , therefore the secondary fields $\theta(t, s)$ are also stationary in space-time.

Next, the computation of the four secondary fields is described, including equations for the pre-transform random fields and the respective transformation functions.

$3.2.1$ $\| U(t,s)$

The pre-transform Gaussian field $U^*(t, s)$ satisfies the basic stochastic model Equation (2):

$$
\frac{\partial U^*}{\partial t} + U_U \frac{\partial U^*}{\partial s} + \rho_U U^* - v_U \frac{\partial^2 U^*}{\partial s^2} = \sigma_U \, \alpha_U(t, s), \quad (18)
$$

where $U_{II}, \rho_{II}, V_{II},$ and σ_{II} are constant and non-random hyperparameters and α_U is the white noise independent of α . The transformation $U^*(t, s) \mapsto U(t, s)$ is simply

$$
U(t,s) = \overline{U} + U^*(t,s),\tag{19}
$$

where \overline{U} is the unperturbed value of U (a non-random scalar).

From linearity of Equations (18) and (19), $U(t, s)$ is a Gaussian random field.

3.2.2 $\sigma(t,s)$

The pre-transform Gaussian field $\sigma^*(t, s)$ satisfies

$$
\frac{\partial \sigma^*}{\partial t} + U_{\sigma} \frac{\partial \sigma^*}{\partial s} + \rho_{\sigma} \sigma^* - \nu_{\sigma} \frac{\partial^2 \sigma^*}{\partial s^2} = \sigma_{\sigma} \alpha_{\sigma}(t, s), \qquad (20)
$$

where U_{σ} , ρ_{σ} , v_{σ} , and σ_{σ} are the hyperparameters and α_{σ} is the independent white noise field.

To define the transformation $\sigma^*(t, s) \mapsto \sigma(t, s)$, we require that the field $\sigma(t, s)$ should be positive and have zero probability density at $\sigma = 0$. To meet these requirements, we postulate that

$$
\sigma(t,s) = \overline{\sigma} \cdot g\{\sigma^*(t,s)\},\tag{21}
$$

where $\overline{\sigma}$ is the unperturbed value of σ and $g(z)$ is the transformation function selected to be the scaled and shifted logistic function (also known as the sigmoid function in machine learning):

$$
g(z) := \frac{1 + e^b}{1 + e^{b-z}},
$$
 (22)

where *b* is the constant. The function $g(z)$ has the following property: it behaves like the ordinary exponential function everywhere except for $z \gg b$, where the exponential growth is tempered (moderated). Indeed, it exponentially decays as $z \rightarrow -\infty$. Like exp(z), it is equal to 1 at $z = 0$. With $b > 0$, $g(z)$ saturates as $z \to \infty$ at the level $1 + e^{b}$; this is the main difference of g from the exponential function and the reason why we replace $exp(z)$ by $g(z)$: to avoid too large values in $\sigma(t, s)$, which can give rise to unrealistically large spikes in ξ . We will refer to *b* as the *g*-function's saturation hyperparameter. For $b = 1$, the function $g(z)$ in plotted in Figure 1 alongside the exponential function.

Due to nonlinearity of the transformation function g , the field $\sigma(t, s)$ is non-Gaussian. Its pointwise distribution is known as logit-normal or logit-Gaussian.

3.2.3 $\int \rho(t,s) \text{ and } \nu(t,s)$

The remaining two secondary fields $\rho(t, s)$ and $v(t, s)$ (denoted here generically by ψ) are modelled in a way similar to $\sigma(t, s)$, the only difference being the transformation

Figure 1 Logistic function $g(z)$ with $b = 1$ and the exponential function

function. Specifically, the pre-transform Gaussian field ψ^* satisfies

$$
\frac{\partial \psi^*}{\partial t} + U_{\psi} \frac{\partial \psi^*}{\partial s} + \rho_{\psi} \psi^* - V_{\psi} \frac{\partial^2 \psi^*}{\partial s^2} = \sigma_{\psi} \alpha_{\psi}(t, s), \quad (23)
$$

where, again, U_{ψ} , ρ_{ψ} , v_{ψ} , and σ_{ψ} , are the hyperparameters and α_{ψ} is the independent white noise field.

The transformation function is defined here to be

$$
\psi(t,s) = \overline{\psi} \cdot \left[(1 + \varepsilon_{\psi}) \cdot g\{\psi^*(t,s)\} - \varepsilon_{\psi} \right],\qquad(24)
$$

where $\overline{\psi}$ is the unperturbed value of ψ , the function g is the same as in Section 3.2.2 and controlled by the same saturation hyperparameter b, and ε_{ψ} is the small non-negative constant. The hyperparameter ε_w is introduced to allow for small negative values of ψ (i.e. of ρ and ν). This will take place if the pre-transform Gaussian field ψ^* happens to take a large negative value. Allowing for negative values of the decay coefficient ρ and the diffusion coefficient ν is motivated by the desire to introduce an intermittent instability into the model.

Like $\sigma(t, s)$, the fields $\rho(t, s)$ and $v(t, s)$ are logit-Gaussian.

3.3 Third level of the hierarchy: hyperparameters and external parameters

In contrast to the popular Lorenz'96 model (Lorenz and Emanuel, 1998), which has only one parameter F , DSADM has many parameters as discussed above (in total, there are 23 hyperparameters). It is convenient to specify the hyperparameters using a set of more sensible *external* parameters as described next.

The advection velocities \overline{U} and U_{θ} (for $\theta = U, \rho, \nu, \sigma$) are specified directly as they have a clear physical meaning.

The unperturbed hyperparameters $\overline{\theta} = \overline{\rho}, \overline{v}, \overline{\sigma}$ are specified from the unperturbed external parameters \overline{L} , \overline{T} , and $\overline{SD(\xi)}$

TABLE 1 Short list of external parameters of DSADM

External parameter	Dependent hyperparameters	Which feature of the model is controlled?	
$SD(\xi)$	$\overline{\sigma}$	Mean $SD(\xi)$ (roughly)	
\overline{U}	$U, U_U, U_\rho, U_\nu, U_\sigma$	All advection velocities	
\overline{L}	$\overline{\rho}, \overline{\nu}$	Mean length-scale of ξ (roughly)	
L^*	$\rho_U, \rho_\rho, \rho_\nu, \rho_\sigma$	Length-scale and time-scale	
	$v_U, v_\rho, v_\nu, v_\sigma$	of non-stationarity	
$SD(U^*)$	σ _U	Strength of non-stationarity	
$x_{\rho}, x_{\nu}, x_{\sigma}$	$\sigma_\rho, \sigma_\nu, \sigma_\sigma$	Strength of non-stationarity	
π_{ρ}, π_{ν}	$\varepsilon_{\rho}, \varepsilon_{\nu}$	Strength of non-stationarity	
		(portions of time ρ and ν are negative)	

using Equations (16), (14), and (12).^{[3](#page-6-0)} The latter determine the desired typical spatial scale, time-scale, and variance of the field $\xi(t, s)$. Note that in the non-stationary regime, \overline{L} , \overline{T} , and $\overline{SD(\xi)}$ are not exactly equal to the mean values of the length-scale, time-scale, and variance of $\xi(t, s)$, respectively, because closed-form expressions for the latter are not available.

Likewise, for any pre-transform Gaussian field θ^* = $U^*, \rho^*, v^*, \sigma^*$, the respective hyperparameters ρ_θ and v_θ are calculated from the external parameters L_{θ} and T_{θ} using Equations (16) and (14). L_{θ} and T_{θ} are the user-specified *spatial and temporal scales of non-stationarity*.

The hyperparameters σ_{θ} are specified as follows. σ_{U} is calculated from the external parameter $SD(U^*)$ using Equation (12). For each of the other three secondary fields, $\theta = \rho, v, \sigma$, we select $x_{\theta} = \exp\{SD(\theta^*)\}$ to be the respective external parameter. This choice is motivated by the fact that these fields are nonlinearly transformed Gaussian fields. As $g(z)$ (defined in Equation (22) and shown in Figure 1) is a "tempered" exponential function, it is worth measuring the standard deviation of, say, the σ^* field on the log scale: $SD(\sigma^*) = \log \chi_{\sigma}$, so that the typical deviation of the transformed field σ from its unperturbed value $\overline{\sigma}$ is x_{σ} times. The external parameters SD(U^*), x_{σ} , x_{ρ} , and x_{ν} determine the *strength* of non-stationarity.

Finally, for the fields $\psi = \rho$, ν , we parametrize ε_{ψ} (which, we recall, controls the occurrence of *negative* values of ψ) using the probability π_{ψ} that $\psi(t, s)$ is negative:

$$
P\{\psi(t,s) < 0\} = \pi_{\psi}.\tag{25}
$$

Substituting ψ from Equation (24) into Equation (25) and utilizing monotonicity of the transformation function g and Gaussianity of the field ψ^* , we easily come up with a relation between π_{ψ} and ε_{ψ} (the elementary formulas are omitted). The external parameters π_{ρ} and π_{v} thus determine how often and how strong *local instabilities* can be. (Actually, π_w

slightly impacts also sup ψ , but this is a minor effect and it can be ignored in this application.)

To facilitate the use of DSADM, we introduce a reduced set of external parameters using the following constraints.

- (i) The unperturbed advection velocity \overline{U} and all advection velocities U_{θ} are equal to each other.
- (ii) The length-scale hyperparameters L_{θ} for all pre-transform fields θ^* are selected to be equal to the common value L^* , the *length-scale of non-stationarity*.
- (iii) The time-scale hyperparameters \overline{T} and T_{θ} are specified to be equal to $\overline{L}/V_{\text{char}}$ and $L_{\theta}/V_{\text{char}} = L^*/V_{\text{char}}$, respectively, where V_{char} is the *characteristic velocity*.
- (iv) The g-function's saturation hyperparameter b is set to 1.

Having selected a meaningful value of V_{char} , the user can control the typical time- and length-scales of ξ , the strength of non-stationarity, and the time- and length-scales of non-stationarity – using the reduced set of the ten external parameters listed in Table 1.

4 PROPERTIES AND CAPABILITIES OF DSADM

4.1 Non-stationarity

Here, we show that, given the secondary fields, the solution to DSADM is indeed a non-stationary in space and time random field. Its spatiotemporal covariances are themselves random (as they depend on the random secondary fields) and appear to be stationary processes in space-time.

Let us rewrite the basic Equation (17) as the linear stochastic state-space model

$$
\frac{\mathrm{d}\xi(t)}{\mathrm{d}t} = \Phi\{\underline{\theta}(t)\}\,\xi(t) + \Sigma\{\underline{\theta}(t)\}\,\alpha(t),\tag{26}
$$

where ξ and θ stand for the vectors of the spatially gridded fields $\xi(t, s)$ and $\theta(t, s)$, respectively, $\alpha(t)$ is the space-discrete and time-continuous white noise, $\Phi(\theta)$ is the spatial operator dependent on the spatial fields θ , and $\Sigma(\theta)$ is the diagonal

³Note that for any field θ , $\overline{\theta}$ is its pointwise *median* (for $\theta = U$ it is also the mean), hence the overbar notation.

matrix whose application to α represents the forcing term $\sigma \alpha$ in Equation (17).

Note that DSADM is intended to be used as a model-of-truth in which the secondary fields, once computed in an experiment, are held fixed. In this default setting, Equation (26) implies that DSADM is a linear non-autonomous stochastic model (because Φ and Σ are explicit functions of time).

Discretizing Equation (26) in time (using an implicit time-differencing scheme) yields the equation for the time-discrete signal ξ :

$$
\xi_k = \mathbf{F}(\underline{\theta}_k) \{ \xi_{k-1} + \Sigma(\underline{\theta}_k) \alpha_k \},\tag{27}
$$

where ξ_k is the field on the spatial grid (an *n*-vector), $k =$ 1*,* 2*,* … labels the time instant, **F** is the model operator, and components of the vector α_k are independent Gaussian variables with mean zero and variance $1/(\Delta s \Delta t)$. Here Δs = $2\pi R/n$ is the spatial grid spacing and Δt the time step. The initial condition ξ_0 has mean zero and the known covariance matrix Γ_0 . Equation (27) allows us to find the conditional covariance matrix of the spatial field ξ_k given the secondary fields $\underline{\theta}_{1:k}$ (the notation 1 : k designates all time instants from 1 to k), that is, $\Gamma_k = \mathbb{E}(\xi_k \xi_k^T | \underline{\theta}_{1:k})$, using the recursion

$$
\Gamma_k = \mathbf{F}(\underline{\theta}_k) \Gamma_{k-1} \mathbf{F}(\underline{\theta}_k)^{\mathrm{T}} + \mathbf{Q}(\underline{\theta}_k),\tag{28}
$$

where $Q = F\Sigma^2 F^T/(\Delta s \Delta t)$ is the covariance matrix of the forcing. The lagged covariances $\Gamma_{kj} = \mathbb{E}(\xi_k \xi_j^T | \underline{\theta}_{1:k})$ (where $j < k$) can be found by right-multiplying Equation (27) by ξ_j^T and taking expectation (note that for $j < k$, ξ_j is independent of α_k :

$$
\Gamma_{kj} = \mathbf{F}(\underline{\theta}_k) \Gamma_{k-1,j}.
$$
 (29)

Now recall that the secondary fields θ are random fields. Then, Equations (28) and (29) imply that the field (signal) covariances Γ_k and Γ_{kj} are (matrix variate) random processes. It can be shown that if $\rho(t, s) > 0$ and $v(t, s) > 0$ (for small enough π _o and π _v this is true most of the time), then **F** (as a matrix operator) is a contraction. Therefore, the dependence of Γ_k and Γ_{kj} on Γ_0 fades out as $k \to \infty$ and only the secondary fields θ determine the covariances of $\xi(t, s)$. Moreover, since the fields θ are stationary in space-time and **F** and **Q** are space- and time-invariant (as operators acting on θ), the field covariances Γ_k and Γ_{kj} are stationary random processes in space-time.

Below, we will be particularly interested in the following two aspects of the spatial covariances Γ_k or, in continuous notation and omitting for brevity the dependence on the secondary fields, $\gamma(t, t; s, s')$:

(a) the time- and space-specific *field variance* $\text{Var} \xi(t, s) =$ $\gamma(t, t; s, s)$ and

(b) a time- and space-specific *spatial scale* defined to be, say, the local macroscale:

$$
\Lambda_{\xi}(t,s) = \frac{1}{2 \operatorname{Var} \xi(t,s)} \int_{\mathbb{S}^1(R)} \gamma(t,t;s,s') \, \mathrm{d}s'. \tag{30}
$$

Then, from stationarity of $\gamma(t, t; s, s')$, it follows that both Var $\xi(t, s)$ and $\Lambda_{\xi}(t, s)$ are stationary in space-time random fields.

4.2 Estimation of true covariances

Instead of the recursive computation of the field covariances using Equation (28), one can *estimate* them as accurately as needed by running DSADM Equation (17) P times with independent realizations of the forcing $\alpha(t, s)$ and with the fields $\theta(t, s)$ held fixed. The spatial covariances $\gamma(t, t; s_1, s_2)$ can then be estimated from the resulting sample $\{\xi^{(p)}(t,s)\}_{p=1}^P$ as the usual sample covariances.

More importantly, this approach can be used to estimate time- and space-specific *true filtering distributions* of any filter in question. Note that these distributions and their parameters (e.g. true background-error covariances) are hard to obtain, even knowing the model of truth. Indeed, on the one hand, a sub-optimal filter, because of its approximate nature, can only yield estimates which are inexact and, likely, biased (hence the need for covariance inflation in EnKF). On the other hand, an exact (e.g. Kalman) filter cannot be used here either – because the error statistics of the approximate filter and the exact filter differ.

To estimate, say, true background-error covariances of a filter, one may use the above fields $\{\xi^{(p)}(t,s)\}_{p=1}^P$ as *F* "truths," generate P sets of "synthetic" observations (with independent errors) in space and time, and perform P assimilation runs. Then, at each time, one may compute P background-error vectors (by subtracting the truth $\xi^{(p)}(t, s)$ from the respective background field) and then, finally, compute their (time-specific) sample covariance matrix. We experimented with DSADM on a 60-point spatial grid and found that the sample size $P = 5,000-10,000$ was enough to accurately estimate true time-specific error covariances (not shown). This approach is similar to that by Bishop and Satterfield (2013); the difference is that in our approach the truth is random whilst they assume that the truth is deterministic.

4.3 Instability

The nonlinear deterministic models mentioned in the Introduction are chaotic, that is, having unstable modes (positive Lyapunov exponents), whereas DSADM is stochastic but experiencing intermittent instabilities due to the possibility for ρ and ν to attain negative values. In the deterministic models instabilities are curbed by nonlinearity, whereas in DSADM, these are limited by the time the random fields ρ and ν remain negative.

4.4 Gaussianity

Given θ , DSADM is linear, hence $\xi(t, s)$ is conditionally Gaussian. Unconditionally, $\xi(t, s)$ is a continuous mixture of zero-mean Gaussian distributions and so has a non-Gaussian distribution with heavy tails⁴.

4.5 Unbeatable benchmark filter

Linearity of DSADM makes it possible to use the exact Kalman filter and thus to know how far from the optimal performance the filter in question is, which is always useful but often not possible with nonlinear deterministic models of truth.

In principle, non-stationarity combined with linearity can be achieved by taking a nonlinear model (like those mentioned in the Introduction) and using its tangent linear model. We have tried that with the Lorenz'96 model and found that even a tiny imposed model-error perturbation leads to an explosive growth of the forecast perturbation. Perhaps some diffusion is needed to stabilize the model. This would complicate the model but, more significantly, the tangent linear model does not allow us to explicitly specify the various characteristics of non-stationarity which the DSADM does.

5 BEHAVIOUR OF THE MODEL

In this section, we numerically examine solutions to DSADM and some aspects of their spatiotemporal distributions.

5.1 Model set-up

The model's differential equations were solved numerically using an implicit upwind finite-difference scheme. The computational grid had $n = 60$ points on the circle of radius $R =$ 6, 370 km. The model integration time step was $\Delta t = 6$ hr. The external parameters of DSADM were selected to resemble the spatiotemporal structure of a mid-tropospheric meteorological field like temperature or geopotential, with one caveat: the time- and length-scales were chosen to be about twice as large as the respective meteorological scales (for the fields to be reasonably resolved on the 60-point spatial grid).

We specified $\overline{U} = 10 \,\text{m s}^{-1}$, $\overline{L} = 5\Delta s = 3,300 \,\text{km}$, L^* = 2L (it is meaningful to assume that the *structural*

change in the field occurs at a larger space- and time-scale than the change in the random field itself), $SD(\xi) = 5$ (selected arbitrarily and does not impact any conclusions), $SD(U^*) = 10 \text{ m s}^{-1}, b = 1, x_{\rho} = x_{\nu} = x_{\sigma} = 3, \pi_{\rho} = 0.02,$ $\pi_v = 0.01$, $V_{\text{char}} = 3 \text{ m s}^{-1}$ (tuned to give rise to the desired mean time-scale).

$5.2 \quad \frac{\xi(t,s)}{\xi(t,s)}$ plots

Figure 2 compares typical spatiotemporal segments of solutions to the stationary stochastic model, Equation (2), (panel a) and DSADM, Equation (17), (panel b). One can see that, indeed, the non-stationary field had different structures at different areas across the domain, whilst the stationary field had the same structure everywhere. In particular, one can spot areas where the non-stationary field experienced more small-scale (large-scale) fluctuation than in the rest of the plot. These spots correspond to small (large) values of the local length-scale Λ_{ξ} (Figure 3b).

5.3 Non-stationarity

Figure 3 shows two characteristics of the non-stationarity pattern: the local log-variance $\log_{10}(\text{Var } \xi(t, s))$ and the local length-scale $\Lambda_{\xi}(t, s)$ (defined in Equation (30)). Both were calculated from the spatial covariance matrix Γ_k computed following Equation (28). We recall that the variability in Var $\xi(t, s)$ and $\Lambda_{\xi}(t, s)$ was induced by the simulated secondary fields $\theta(t, s)$, on which \mathbf{F}_k and \mathbf{Q}_k depend. The same realizations of the secondary fields were used to plot both Figure 3 and Figure 2b. In Figure 3, one can see the substantial degree of non-stationarity (note that in the stationary case both Var ξ and Λ_{ξ} are constant). Specifically, in Figure 3a, the ratio of the maximum to the minimum field variance Var $\xi(t, s)$ is seen to be greater than two orders of magnitude. The same ratio for the length-scale $\Lambda_{\varepsilon}(t, s)$ (Figure 3b) was about 5, which also indicates a significant degree of variation.

Figure 4 displays how diverse the true spatial *correlations* were. Note that Figure 3a illustrates the non-stationarity of the field's magnitude, whereas Figures 3b and 4 highlight the non-stationarity of the field's spatial structure.

Thus, DSADM is capable of generating significantly non-stationary random fields.

5.4 Gaussianity

As noted above, conditionally on the secondary fields, $\xi(t, s)$ is Gaussian by construction. This is the way DSADM is used in this study. In principle, it can be used for other purposes without conditioning on the secondary fields, so that in each model run both the forcing and the model coefficients are random (and the model becomes autonomous). In this setting, the generated field ξ is no longer Gaussian. Numerical

⁴ Indeed, one can prove using the Jensen inequality that the *kurtosis* of a non-degenerate mixture of this kind (a scale mixture) is always greater than 3 (the Gaussian kurtosis). A positive excess kurtosis means, normally, more probability mass in the tails of the distribution than in the tails of the Gaussian distribution with the same mean and variance.

Figure 2 Realizations of $\xi(t, s)$ generated by (a) the stationary model Equation (2), and (b) the non-stationary DSADM Equation (17)

Figure 3 Space-time plots of estimated (a) field log-variance \log_{10} {Var $\xi(t, s)$ }, and (b) spatial macroscale $\Lambda_{\xi}(t, s)$

experiments confirmed that the unconditional distribution of $\xi(t, s)$ was indeed non-Gaussian with heavy tails (as we anticipated in Section 4.4). The non-Gaussianity was stronger for larger magnitudes of the secondary fields (not shown).

5.5 Forward and inverse modelling

Building DSADM can be called *forward* hierarchical modelling. That is, we have formulated a (hopefully) reasonable hierarchical model and an algorithm to compute realizations of the first-level random field $\xi(t, s)$ given the third-level hyperparameters ϕ . A harder problem is *inverse* modelling, that is, the inference about the model parameters – in our case the parameter fields $\theta(t, s)$ – from a number of realizations (an ensemble) of the field ξ . This is the classical hierarchical Bayesian problem, which is beyond the scope of this study but can be relevant in a broader context of non-stationary spatial and spatiotemporal field modelling.

Figure 4 Spatial correlations of the field ξ with respect to 30 randomly selected points in space-time

6 HYBRID HBEF (HHBEF) FILTER

In the rest of the paper, we use DSADM to study the impact of non-stationarity on the performance of hybrid ensemble filters. We formulate a new filter that extends HBEF (in which ensemble covariances are blended with recent-past time-smoothed covariances, see Tsyrulnikov and Rakitko, 2017) by introducing two additional kinds of hybridization mentioned in the Introduction: blending with climatological covariances (as in EnVar) and blending with spatially neighbouring covariances (spatial smoothing). We start with a review of HBEF and then introduce HHBEF.

6.1 HBEF

Tsyrulnikov and Rakitko (2017) developed, building on Myrseth and Omre (2010), Bocquet (2011), and Bocquet *et al.* (2015), their Hierarchical Bayes ensemble Kalman filter (HBEF), which cyclically updates prior covariances using ensemble members as generalized observations and the estimated previous-cycle covariances as a background. The (Bayesian) update relies on the inverse-Wishart matrix-variate hyperprior distribution for the unknown, and thus assumed random, covariance matrices.

HBEF updates the model-error and the predictability-error covariance matrices, but here we consider a simplified design in which the background-error covariance matrix **is cyclically updated. In the simplest version of HBEF,** the mean-square-optimal posterior (analysis) estimate \mathbf{B}_k^a of the unknown true matrix \mathbf{B}_k is the linear combination of the *background* \mathbf{B}_k^f and the (localized and inflated) sample covariance matrix, \mathbf{B}_k^e :

$$
\mathbf{B}_{k}^{\mathrm{a}} = \frac{\partial \mathbf{B}_{k}^{\mathrm{f}} + N \mathbf{B}_{k}^{\mathrm{e}}}{\partial + N},\tag{31}
$$

where N is the ensemble size and $\theta > 0$ the so-called sharpness parameter of the inverse Wishart distribution (defined in appendix A of Tsyrulnikov and Rakitko, 2017). The background \mathbf{B}_k^f is provided by the persistence forecast:

$$
\mathbf{B}_k^{\text{f}} = \mathbf{B}_{k-1}^{\text{a}}.\tag{32}
$$

The matrix \mathbf{B}_k^a computed according to Equation (31) is used in the analysis of the state ξ_k as the background-error covariance matrix. The posterior (analysis) ensemble is generated in HBEF (and in all other filters in this study) in the same way as in the stochastic EnKF (e.g. Houtekamer and Zhang, 2016).

From Equations (31) and (32), it follows that \mathbf{B}_k^a satisfies the first-order autoregressive equation

$$
\mathbf{B}_{k}^{\text{a}} = \mu \mathbf{B}_{k-1}^{\text{a}} + (1 - \mu) \mathbf{B}_{k}^{\text{e}},\tag{33}
$$

where $\mu = \theta/(\theta + N) < 1$. Equation (33) implies a kind of time smoothing of ensemble covariances. In particular, $\mathbf{B}_{k-1}^{\text{a}}$ is the time-smoothed recent-past ensemble covariance matrix denoted in what follows by \mathbf{B}_k^r . With this notation, Equation (33) shows that \mathbf{B}_k^a is a (convex) linear combination of \mathbf{B}_k^r and \mathbf{B}_k^e . The role of \mathbf{B}_k^r in HBEF is two-fold. First, if spatial non-stationarity has some memory (which is highly likely in realistic systems), then \mathbf{B}_k^r brings this past memory to the current assimilation cycle, improving the accuracy of the resulting estimate of the true background-error covariance matrix \mathbf{B}_k . Second, due to time smoothing (i.e. averaging), the sampling noise is reduced, leading to a more accurate estimate \mathbf{B}_k^a .

6.2 HHBEF: blending with static covariances

The first idea of the new hybrid-HBEF (HHBEF) filter is to replace the HBEF's persistence forecast for the covariance matrix \mathbf{B}_k (Equation (32)) with a regression-to-the-mean forecast:

$$
\mathbf{B}_k^{\mathrm{f}} = w \mathbf{B}_{k-1}^{\mathrm{a}} + (1 - w) \mathbf{B}^{\mathrm{c}}.
$$
 (34)

Here, \mathbf{B}^c is the climatological covariance matrix and $w \in$ [0*,* 1] is the scalar weight.

6.3 Spatial smoothing of the covariances

The second idea of HHBEF is to accommodate spatially smoothed ensemble covariances. Here, we review the approach by Buehner and Charron (2007), who studied spectral-space localization, found it useful in reducing sampling noise, and noted that it is equivalent to a spatial smoothing of the covariance function. Their equation 9, is rewritten here as

$$
\check{B}(s_1, s_2) = \int_{\mathbb{S}^1(R)} \kappa(s) \, B(s_1 - s, s_2 - s) \, \mathrm{d}s,\tag{35}
$$

where the check $\check{\ }$ designates the smoothing and $\kappa(s) \geq 0$ is the weighting (averaging) function. In the space-discrete case, Equation (35) can be approximated as follows:

$$
B_{ii'} = \sum_{s=-n/2+1}^{n/2} \kappa_s B_{i-s,i'-s},
$$
 (36)

where $\kappa_s \geq 0$ are the weights such that $\sum \kappa_s = 1$. In the matrix form Equation (36) can be written as

$$
\check{\mathbf{B}} = \sum_{s=-n/2+1}^{n/2} \kappa_s \mathcal{F}^s \mathbf{B} \mathcal{F}^{-s},\tag{37}
$$

where $\mathcal F$ is the forward-shift operator. From Equation (37) with non-negative κ_s and given that $\mathcal{F}^{-1} = \mathcal{F}^T$, it immediately follows that the smoothing is properly defined in the sense that the resulting smoothed covariance matrix $\tilde{\mathbf{B}}$ is non-negatively definite.

If **B** = $\frac{1}{N-1} \sum_{i} e_i e_i^T$, where e_i is the *l*th ensemble perturbation, $l = 1, ..., N$, then Equation (37) implies that the spatial smoothing of the sample covariance matrix can be performed by spatially shifting (through the repeated application of the forward and backward shift operators) and weighting (through the multiplication by $\sqrt{\kappa_s}$) ensemble perturbations – as was proposed and tested by Buehner and Charron (2007).

6.4 HHBEF's hybrid covariances

Combining Equations (34) with (31), where B^e is replaced by its spatially smoothed version $\tilde{\mathbf{B}}^e$ computed according to Equation (37), yields the HHBEF's analogue of the HBEF's Equation (33):

$$
\mathbf{B}_{k}^{\text{a}} = \mu w \mathbf{B}_{k-1}^{\text{a}} + (1 - \mu) \check{\mathbf{B}}_{k}^{\text{e}} + \mu (1 - w) \mathbf{B}^{\text{c}}, \qquad (38)
$$

where μ is the weight of \mathbf{B}_k^f relative to $\check{\mathbf{B}}_k^e$. Solving Equation (38) (which is a forced linear difference equation; the derivation is omitted) shows that, for $\mu w < 1$ and after an initial transient, \mathbf{B}_k^a has the four components:

$$
\mathbf{B}_k^a = w_e \mathbf{B}_k^e + w_{es} \mathbf{B}_k^{es} + w_r \mathbf{B}_k^r + w_c \mathbf{B}^c, \tag{39}
$$

where $w_e = (1 - \mu)\kappa_0$ is the weight of the current ensemble covariances \mathbf{B}_k^e ,

 $w_{\rm es} = (1 - \mu)(1 - \kappa_0)$ is the weight of the *spatially smoothed* current ensemble covariances $\mathbf{B}_k^{\text{es}} = \check{\mathbf{B}}_k^{\text{e}} - \dot{\mathbf{B}}_k^{\text{e}}$

 $w_c = \mu(1 - w)/(1 - \mu w)$ is the weight of the climatological covariances **B**^c, and

 $w_r = \mu w (1 - \mu) / (1 - \mu w)$ is the weight of the spatiotemporally smoothed recent-past ensemble covariances \mathbf{B}_k^r :

$$
\mathbf{B}_{k}^{\mathrm{r}} = \frac{\check{\mathbf{B}}_{k-1}^{\mathrm{e}} + \mu w \check{\mathbf{B}}_{k-2}^{\mathrm{e}} + (\mu w)^{2} \check{\mathbf{B}}_{k-3}^{\mathrm{e}} + \dots}{1 + \mu w + (\mu w)^{2} + \dots}.
$$
 (40)

The smoothing time-scale (measured in assimilation cycles) is $-1/\log(uw)$.

Equations (37) and (38) show that $w = 1$ and $\kappa_s = \delta_{s0}$ (where δ_{s0} is the Kronecker delta) reduces HHBEF to HBEF, $w = 0$ to EnVar, $\mu = 0$ and $\kappa_s = \delta_{s0}$ to pure EnKF, and $w = 0$, $\mu = 1$ recovers the filter with static background-error covariances.

Thus, HHBEF employs blending of (localized) sample covariances with both climatological and spatiotemporally smoothed background-error covariances.

7 PERFORMANCE OF THE THREE COVARIANCE-BLENDING TECHNIQUES UNDER NON-STATIONARITY

In this section, we experimentally study how temporal smoothing of background-error covariances, their spatial smoothing, and their mixing with climatology affects the performance of the classical stochastic EnKF (e.g. Houtekamer and Zhang, 2016) and HHBEF under different regimes of non-stationarity.

7.1 "Twin" experimental methodology

In the experiments below, the truth was generated using the discretized Equation (17), that is, Equation (27): ξ_k = $\mathbf{F}_k \boldsymbol{\xi}_{k-1} + \boldsymbol{\varepsilon}_k$, where $\boldsymbol{\varepsilon}_k$ is the time and space-discrete model forcing. The forecast operator \mathbf{F}_k was available to the filters, so that, given the analysis ξ_{k-1}^a , the next-time-instant forecast was $\xi_k^f = \mathbf{F}_k \xi_{k-1}^a$. In this setting, the model forcing ε_k becomes the model error (e.g. Tsyrulnikov and Gayfulin (2017)), with its covariance matrix \mathbf{Q}_k also available to the filters. We reiterate that in each experiment, the secondary fields (which determine the forecast operator and the model-error statistics), once generated, were kept fixed.

The above setting implies that *non-identical* true and model twins were used, with the only difference between the two models being the stochastic model error. In contrast to using *different models* to generate the truth and to perform the forecast, this approach provides us with the exact knowledge (and in principle, full control) of the model error statistics. It is this knowledge (along with linearity) that justifies the use of the exact Kalman filter.

7.2 Data assimilation set-up

The default DSADM set-up was the same as in Section 5.1. The default ensemble size was $N = 10$. The set-up of the filters and the observation network were chosen with the intention to mimic – as much as possible with a 1D model – the set-up of a realistic meteorological data assimilation scheme.

The time interval between consecutive analyses, T_a , was selected using the following criterion. The DSADM field correlation with lag T_a should approach the typical meteorological-field correlation at lag 6 hr (the most widely used in operational practice assimilation cycle in global schemes). On the one hand, the mid-tropospheric 6 hr time correlation was estimated by Seaman (1975, figure 1b) using radiosonde data to be about 0.8 for the wind fields and can be interpolated to the value of about 0.95 for geopotential (Olevskaya, 1968) (that is, about 0.9 on average). On the other hand, this 0.9 correlation in the DSADM field (with the default set-up) occurred, on average, at the 12 hr lag. So, we set $T_a = 12$ hr. This is twice the typical assimilation cycle in real-world systems and is consistent with the selection of L and \overline{T} in Section 5.1 (about twice as large as the respective typical scales of meteorological fields).

Observations were generated every assimilation cycle at every tenth spatial grid point. The observation error standard deviation (=6) was selected to ensure that the mean reduction of the forecast-error variance in the analysis was close to the value of 10% reported by Errico and Privé (2014, section 8) for a realistic data assimilation system.

The weights of the spatial covariance smoothing κ_s (Equation (37)) were specified to be a triangular function of the spatial shift s, so that κ_s had a maximum at $s = 0$ and vanished for $|s| > s_{\text{max}}$. Thus, the spatial smoothing was controlled by the single parameter s_{max} (the length-scale of the spatial smoothing measured in grid spacings).

The covariance localization was performed using the popular function by Gaspari and Cohn (1999, equation 4.10).

7.3 Set-up of experiments

The relative roles of the three covariance-blending devices were studied by measuring, first, the performance of EnKF with each device switched on (in turn) versus the pure EnKF, and second, by measuring the performance of HHBEF with each device switched off (again, in turn) versus the full HHBEF. Switching on (off) the mixing with climatology is denoted by $+C$ ($-C$). Similarly, switching on (off) the spatial smoothing of background-error covariances is denoted by $+S$ $(-S)$ and the temporal smoothing by $+T$ ($-T$). Technically, all filters were built within HHBEF.

Each filtering configuration was defined by the triple of the covariance-blending parameters w , μ , and s_{max} as

TABLE 2 Configurations of filters

Filter	w	μ	$S_{\rm max}$	
EnKF	No effect	0	θ	
$EnKF + C (EnVar)$	Ω	Tuned	Ω	
$EnKF + S$	No effect	0	Tuned	
$EnKF + T (HBEF)$	1	Tuned	Ω	
HHBEF	Tuned	Tuned	Tuned	
$H H BEF - C$	1	Tuned	Tuned	
$HHBEF-S$	Tuned	Tuned	0	
HHBEF- T		Tuned	Tuned	

indicated in Table 2. In each experiment, the free parameters marked in the respective row of Table 2 as "Tuned" as well as the multiplicative covariance inflation and the length-scale of the covariance localization were all simultaneously manually tuned to get the best performance. The tuning was sometimes tedious but the optimum was always well-defined and there was no indication that it was not unique. The search for the optimum was facilitated by the observation that the better the performance of a filter, the less need for any covariance-regularization device.

The performance of a filter, f, was measured by its background root-mean-square error (brmse, with respect to the known truth, in runs with 5,000 assimilation cycles) relative to that of the exact Kalman filter (KF):

$$
rel.err(f) = \frac{brmse(f) - brmse(KF)}{brmse(KF)}.
$$
 (41)

The climatological background-error covariance matrix **B**^c was specified in each experiment to be the time-mean KF's background-error covariance matrix (in a run with 50,000 assimilation cycles) – for simplicity and in an attempt to reproduce the realistic situation in which only a proxy to **B**^c is available.

7.4 Results

As noted in Section 3.3, the non-stationarity grows with the increasing external parameters $SD(U^*)$, x_0 , x_v , x_σ , π_v , π_v , Here we compare the filters in four regimes with different strengths of non-stationarity as detailed in Table 3.

Figure 5 shows the performance scores (defined in Equation (41)) for the four configurations indicated in the upper half of Table 2 plus the full HHBEF. One can see that all three covariance-blending techniques were quite successful in improving the performance of EnKF. This could be expected because the ensemble size $N = 10$ was rather small.

Static covariances were most useful in the low non-stationarity regimes (non-stationarity strengths 0 and 1) while being least useful under strong non-stationarity. This is reasonable because, under stationarity, static covariances suffice for exact filtering, whilst they become useless when

TABLE 3 External parameters of DSADM in the four regimes of non-stationarity

No.	Regime	$SD(U^*)$	$x_0 = x_v = x_\sigma$	$\pi_{\scriptscriptstyle{D}}$	π_{ν}
	Stationary				
	Weakly non-stationary			0.01	θ
	Default non-stationary			0.02	0.01
	Strongly non-stationary	20		0.04	0.02

Figure 5 Performance scores (relative to KF, the lower the better) of the pure EnKF (the upper curve with \times), EnKF with one of the three covariance-blending devices switched on (the three middle curves), and HHBEF (the lowest curve with ∗). The filter notation follows Table 2. The numbers on the x -axis correspond to the numbers in the first column of Table 3 [Colour figure can be viewed at [wileyonlinelibrary.com\]](http://wileyonlinelibrary.com).

the filtering-error statistics are, typically, very different from climatology, which is the case in a highly non-stationary regime.

The temporal covariance smoothing is seen to be beneficial in all regimes, especially under strong non-stationarity, and always more efficient than the spatial covariance smoothing. The reason for the observed systematic advantage of the temporal covariance smoothing over the spatial covariance smoothing is unclear. We may conjecture that neighbouring-in-space covariances bring less information because they are part of the same sample covariance matrix, whereas neighbouring-in-time covariances come from a different sample covariance matrix and can therefore be more independent.

It is seen in Figure 5 that HHBEF was more accurate than the other filters.

To check if the three covariance-blending techniques remained beneficial when used in combination, the four configurations indicated in the lower half of Table 2 were compared: HHBEF versus HHBEF with one of the three covariance-blending devices withheld. Figure 6 displays the results, which were, generally, consistent with those depicted

0.030 HHBEF
HHBEF 0.000 0.010 0.020 0.030 $-C$ HHBEF -S Relative performance score ┻ HHBEF-T Relative performance score HHBEF − T 0.020 0.010 0.000 $\overline{2}$ 3 S trength of non-stationarity

 $S₅$ **Figure 6** As Figure 5, but for the full HHBEF and HHBEF with one of the three covariance-blending devices switched off [Colour

figure can be viewed at [wileyonlinelibrary.com\]](http://wileyonlinelibrary.com).

in Figure 5. Under non-stationarity, all three blending devices remained useful if used jointly except for the strongly non-stationary regime 3, where only time smoothing was beneficial. Not using static covariances led to the worst performance under stationarity and low non-stationarity. Switching off the time smoothing was most detrimental for the default and strong non-stationarity.

It is also of interest to inspect the optimal HHBEF's weights: w_e , w_{es} , w_r , and w_c (Equation (39) and the subsequent paragraph). The weights are plotted in Figure 7 for the same four strengths of non-stationarity as above. Overall, we see that the two dominant sources of hybrid covariances were static covariances (prevailing in the low non-stationarity regimes) and spatiotemporally smoothed covariances (most important if the non-stationarity was high). The role of non-smoothed ensemble covariances increased with the growing non-stationarity. This is meaningful because any covariance blending reduces sampling noise but distorts the flow-dependent signal carried by the non-smoothed ensemble covariances, and this distortion grows when the covariances become more variable. The spatially averaged covariances were beneficial only for non-stationarity strengths 1 and 2 and useless under the strong non-stationarity (and of course under stationarity). This is consistent with Figure 6.

Figure 7 Optimal blending weights of climatological covariances, w_c , non-smoothed current ensemble covariances, w_e , spatially smoothed current ensemble covariances, w_{es} , and spatiotemporally smoothed recent-past covariances, w_r . The numbers on the x-axis correspond to the numbers in the first column of Table 3

We also studied how the effects of the three covariance-blending techniques depend on the *time- and length-scales* of non-stationarity (determined by the external parameter L^* ; Table 1). We found that the effects were quite stable, while blending with spatially and spatiotemporally smoothed covariances was more beneficial for larger L^* (not shown). This can be expected because space and time smoothing is worthwhile only if the covariances vary smoothly in space and time.

Finally, we explored the performance of the three covariance-blending techniques for large ensembles ($N =$ 80 and $N = 300$) under the default non-stationarity. First, we found that for those ensembles, some covariance regularization was still useful. This is reasonable because for the non-regularized sample covariance matrix to be a good estimator of the true covariance matrix, the sample size must be significantly greater than the dimensionality of the matrix, e.g. Bai and Shi (2011, section 2), i.e. $N \gg n$, where, we recall, $n = 60$ in our experiments. Second, the effects of all five regularization techniques (the three blending devices plus covariance localization and multiplicative covariance inflation) were significantly reduced compared with the ensemble sizes $N = 10$ (see above) and $N = 30$ (not shown), which is, of course, meaningful. Third, the effect of blending with climatology was negligibly small (because two factors – variability in the covariances and large ensemble size – both reduce the usefulness of static covariances). Fourth, the effects of time and space smoothing remained positive, with time smoothing being still more useful. Fifth, only with huge ensembles ($N = 3,000$), did EnKF need no more covariance regularization (not shown).

8 CONCLUSIONS

In this paper we have presented a new doubly stochastic advection-diffusion-decay model (DSADM) on the circle. Double stochasticity means that, not only is the model forcing stochastic, but the model coefficients (parameters) are random as well. The parameters are specified to be transformed Gaussian random fields, with each Gaussian field satisfying its own stochastic advection-diffusion-decay model with constant coefficients. Thus, DSADM is hierarchical, built of linear stochastic partial differential equations at two levels in the hierarchy. DSADM is designed to be used as a toy model of truth to test and develop data assimilation methodologies.

The main advantage of DSADM is its capability of generating spatiotemporal random fields with tunable non-stationarity in space and time, while maintaining linearity and Gaussianity. This allows one first, to separate effects of non-stationarity from effects of nonlinearity and non-Gaussianity (which is impossible with nonlinear models of truth). Second, linearity and Gaussianity allow the use of the Kalman filter as an unbeatable benchmark, which, again, is rarely possible with nonlinear models. With a small modification, DSADM can also be used to study the role of non-Gaussianity not caused by nonlinearity (say, non-Gaussianity of observation or model errors).

We have used DSADM to study the impact of non-stationarity on the performance of the following three covariance regularization techniques in EnKF: blending with static, time-smoothed, and space-smoothed background-error covariances. DSADM and the synthetic observation network were set up to resemble (as far as possible with a one-dimensional model) a realistic meteorological data assimilation system. We found that blending with static covariances was most beneficial in filtering regimes with low non-stationarity, while time-smoothing was most useful under medium and high non-stationarity. Space-smoothing was less efficient than time-smoothing. These findings were valid in a wide range of ensemble sizes. The role of time and space smoothing was found larger when the length- and time-scales of non-stationarity patterns were larger than the respective scales of the background-error field itself. A new filter termed HHBEF (Hybrid Hierarchical Bayes Ensemble Filter) which combines all three covariance-blending techniques proved to be most accurate among all configurations of the filters tested.

We believe that these results are relevant for real-world applications, but of course they are model-dependent and the degree of relevance remains to be verified, especially because the model is new.

Thus, the optimal blend of ensemble covariances with climatology as well as with time-smoothed and space-smoothed background-error covariances is found to strongly depend on characteristics of non-stationarity. How large is the actual non-stationarity of the spatiotemporal background-error field in practical data assimilation systems, how large are the timeand space-scales of non-stationarity patterns compared to the respective scales of the background-error field itself, how non-stationarity depends on the weather situation, season, scale, altitude, meteorological field, observation density and accuracy – all these questions remain open. Addressing these questions may help optimize hybrid filters that accommodate various covariance regularization techniques.

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REFERENCES

- Apte, A., Auroux, D. and Ramaswamy, M. (2010) Variational data assimilation for discrete Burgers equation. *Electronic Journal of Differential Equations*, 19, 15–30.
- Arnold, L. (1974) *Stochastic Differential Equations: Theory and applications*. Wiley, New York, NY.
- Bai, J. and Shi, S. (2011) Estimating high-dimensional covariance matrices and its applications. *Annals of Economics and Finance*, 12(2), 199–215.
- Banerjee, S., Carlin, B.P. and Gelfand, A.E. (2015) *Hierarchical Modeling and Analysis for Spatial Data*. CRC Press/Chapman & Hall, Boca Raton, FL.
- Berre, L., Varella, H. and Desroziers, G. (2015) Modelling of flow-dependent ensemble-based background-error correlations using a wavelet formulation in 4D-Var at Météo-France. *Quarterly Journal of the Royal Meteorological Society*, 141(692), 2803–2812.
- Bishop, C.H. and Satterfield, E.A. (2013) Hidden error variance theory. Part I: exposition and analytic model. *Monthly Weather Review*, 141(5), 1454–1468.
- Bocquet, M. (2011) Ensemble Kalman filtering without the intrinsic need for inflation. *Nonlinear Processes in Geophysics*, 18(5), 735–750.
- Bocquet, M., Raanes, P.N. and Hannart, A. (2015) Expanding the validity of the ensemble Kalman filter without the intrinsic need for inflation. *Nonlinear Processes in Geophysics*, 22(6), 645–662.
- Bonavita, M., Hólm, E.V., Isaksen, L. and Fisher, M. (2016) The evolution of the ECMWF hybrid data assimilation system. *Quarterly Journal of the Royal Meteorological Society*, 142, 287–303.
- Buehner, M. and Charron, M. (2007) Spectral and spatial localization of background-error correlations for data assimilation. *Quarterly Journal of the Royal Meteorological Society*, 133, 615–630.
- Buehner, M., Morneau, J. and Charette, C. (2013) Four-dimensional ensemble-variational data assimilation for global deterministic weather prediction. *Nonlinear Processes in Geophysics*, 20(5), 669–682.
- Chatfield, C. (2016) *The Analysis of Time Series: An Introduction*. Chapman & Hall/CRC, Boca Raton, FL.
- Dunlop, M.M., Girolami, M.A., Stuart, A.M. and Teckentrup, A.L. (2018) How deep are deep Gaussian processes? *Journal of Machine Learning Research*, 19(1), 2100–2145.
- Errico, R.M. and Privé, N.C. (2014) An estimate of some analysis-error statistics using the Global Modeling and Assimilation Office observing-system simulation framework. *Quarterly Journal of the Royal Meteorological Society*, 140, 1005–1012.
- Gaspari, G. and Cohn, S.E. (1999) Construction of correlation functions in two and three dimensions. *Quarterly Journal of the Royal Meteorological Society*, 125, 723–757.
- Gustafsson, N., Bojarova, J. and Vignes, O. (2014) A hybrid variational ensemble data assimilation for the HIgh-Resolution Limited-Area Model (HIRLAM). *Nonlinear Processes in Geophysics*, 21(1), 303–323.
- Hamill, T.M. and Snyder, C. (2000) A hybrid ensemble Kalman filter – 3D variational analysis scheme. *Monthly Weather Review*, 128(8), 2905–2919.
- Hansen, J.A. and Smith, L.A. (2001) Probabilistic noise reduction. *Tellus A*, 53(5), 585–598.
- Houtekamer, P. and Zhang, F. (2016) Review of the ensemble Kalman filter for atmospheric data assimilation. *Monthly Weather Review*, 144(12), 4489–4532.
- Kumar, P.R. and Varaiya, P. (2015) *Stochastic Systems: Estimation, Identification, and Adaptive Control*. SIAM, Philadelphia, PA.
- Ledoit, O. and Wolf, M. (2004) A well-conditioned estimator for large-dimensional covariance matrices. *Journal of Multivariate Analysis*, 88(2), 365–411.
- Lindgren, F., Rue, H. and Lindström, J. (2011) An explicit link between Gaussian fields and Gaussian Markov random fields: the stochastic partial differential equation approach. *Journal of the Royal Statistical Society. Series B*, 73(4), 423–498.
- Lorenc, A.C. (2003) The potential of the ensemble Kalman filter for NWP – a comparison with 4D-Var. *Quarterly Journal of the Royal Meteorological Society*, 129, 3183–3203.
- Lorenc, A.C. (2017) Improving ensemble covariances in hybrid variational data assimilation without increasing ensemble size. *Quarterly Journal of the Royal Meteorological Society*, 143, 1062–1072.
- Lorenz, E.N. (1963) Deterministic nonperiodic flow. *Journal of Atmospheric Sciences*, 20(2), 130–141.
- Lorenz, E.N. (2005) Designing chaotic models. *Journal of Atmospheric Sciences*, 62(5), 1574–1587.
- Lorenz, E.N. and Emanuel, K.A. (1998) Optimal sites for supplementary weather observations: simulation with a small model. *Journal of Atmospheric Sciences*, 55(3), 399–414.
- Ménétrier, B. and Auligné, T. (2015) Optimized localization and hybridization to filter ensemble-based covariances. *Monthly Weather Review*, 143(10), 3931–3947.
- Ménétrier, B., Montmerle, T., Michel, Y. and Berre, L. (2015) Linear filtering of sample covariances for ensemble-based data assimilation. Part I: optimality criteria and application to variance filtering and covariance localization. *Monthly Weather Review*, 143(5), 1622–1643.
- Miller, R.N., Ghil, M. and Gauthiez, F. (1994) Advanced data assimilation in strongly nonlinear dynamical systems. *Journal of Atmospheric Sciences*, 51(8), 1037–1056.
- Muccino, J. and Bennett, A. (2002) Generalized inversion of the Korteweg–de Vries equation. *Dynamics of Atmospheres and Oceans*, 35(3), 227–263.
- Myrseth, I. and Omre, H. (2010) Hierarchical ensemble Kalman filter. *SPE Journal*, 15(2), 569–580.
- Olevskaya, S.M. (1968) The use of the canonical-correlation method in analyzing the geopotential field. *Izvestiya Atmospheric and Oceanic Physics*, 4, 1149–1159.
- Piterbarg, L. and Ostrovskii, A. (2013) *Advection and Diffusion in Random Media: Implications for sea surface temperature anomalies*. Springer, Berlin.
- Roininen, L., Girolami, M., Lasanen, S. and Markkanen, M. (2019) Hyperpriors for Matérn fields with applications in Bayesian inversion. *AIMS Inverse Problems & Imaging*, 13, 1–29. [https://doi.org/](https://doi.org/10.3934/ipi.2019001) [10.3934/ipi.2019001.](https://doi.org/10.3934/ipi.2019001)
- Seaman, R. (1975) Distance–time autocorrelation functions for winds in the Australian region. *Australian Meteorological Magazine*, 23(2), 27–40.
- Sigrist, F., Künsch, H.R. and Stahel, W.A. (2015) Stochastic partial differential equation based modelling of large space-time data sets. *Journal of the Royal Statistical Society. Series B*, 77(1), 3–33.
- Tjøstheim, D. (1986) Some doubly stochastic time series models. *Journal of Time Series Analysis*, 7(1), 51–72.
- Tsyroulnikov, M.D. (2001) Proportionality of scales: An isotropy-like property of geophysical fields. *Quarterly Journal of the Royal Meteorological Society*, 127, 2741–2760.
- Tsyrulnikov, M. and Gayfulin, D. (2016) *A spatio-temporal stochastic pattern generator for simulation of uncertainties in geophysical ensemble prediction and ensemble data assimilation*. arXiv:1605.02018
- Tsyrulnikov, M. and Gayfulin, D. (2017) A limited-area spatio-temporal stochastic pattern generator for simulation of uncertainties in ensemble applications. *Meteorologische Zeitschrift*, 26(5), 549–566.
- Tsyrulnikov, M. and Rakitko, A. (2017) A hierarchical Bayes ensemble Kalman filter. *Physica D*, 338, 1–16.
- Whittle, P. (1986) *Systems in Stochastic Equilibrium*. John Wiley & Sons, New York, NY.
- Wikle, C.K., Berliner, L.M. and Cressie, N. (1998) Hierarchical Bayesian space-time models. *Environmental and Ecological Statistics*, 5(2), 117–154.
- Yaglom, A.M. (1987) *Correlation Theory of Stationary and Related Random Functions, Volume 1: Basic Results*. Springer, Berlin.

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